

Calculating subsurface nucleonic production of noble gas nuclides: implications on crustal and mantle K, Th, U abundances



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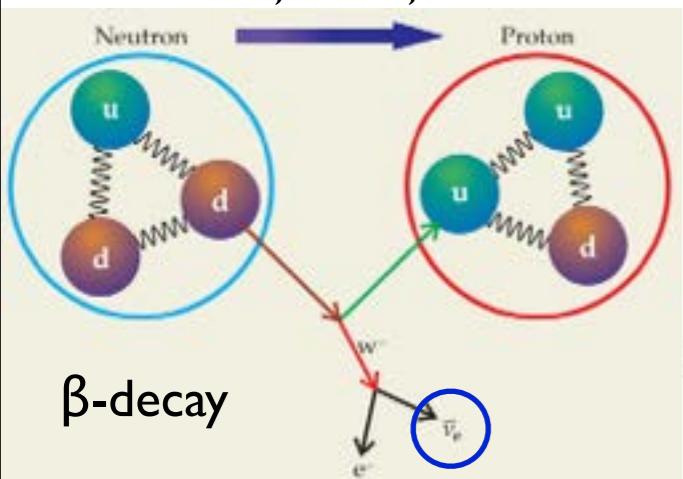
Collaboration with
Bill McDonough (Univ. Maryland)
Sujoy Mukhopadhyay (Harvard)
Lauren Stevens (Univ. Maryland)
Jerry Peterson (CU Boulder)

How much radiogenic heating in the mantle??

What is the Earth made of??

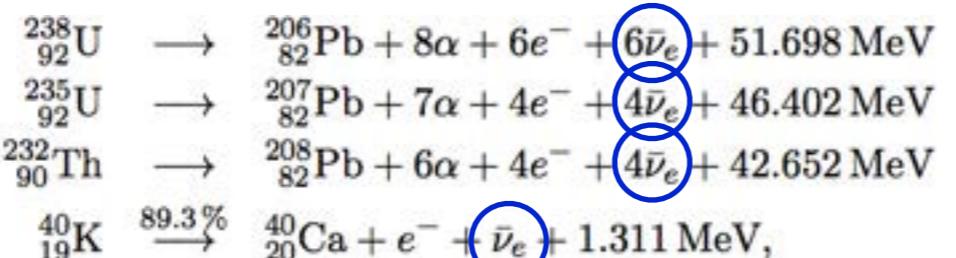
Chemical reservoirs in the mantle??

^{238}U , ^{232}Th , ^{40}K



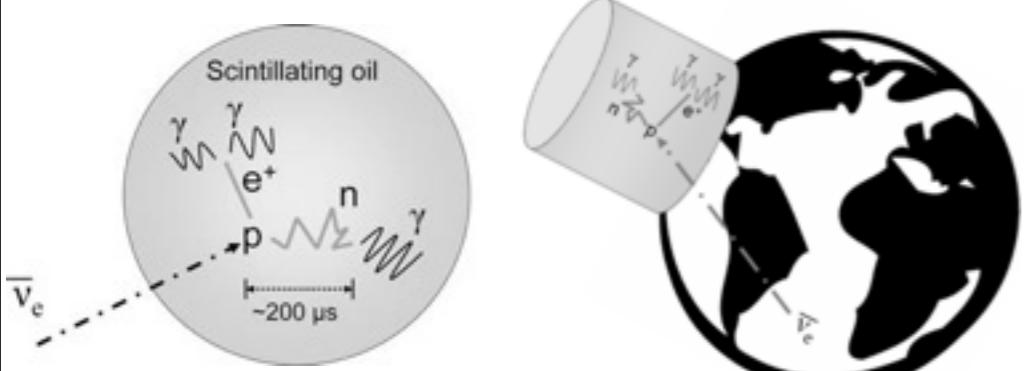
Geoneutrinos:

electron antineutrinos emitted in β -decays of natural radionuclides

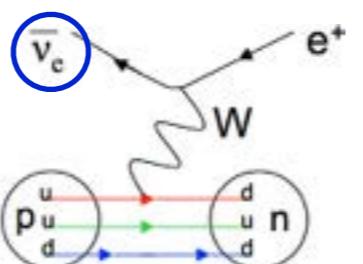


Geo- ν detection possible:

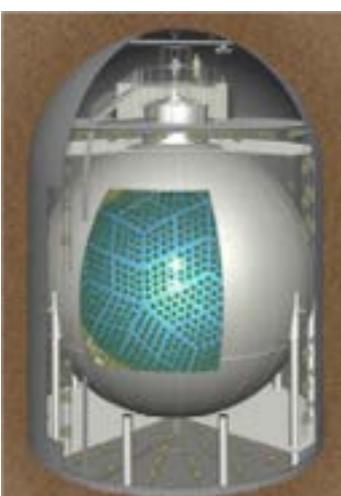
Large liquid scintillator detectors



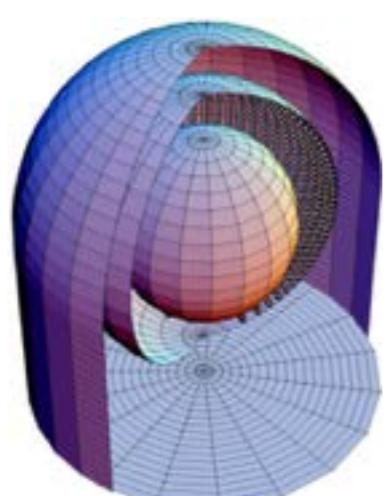
Direct information on
Earth's deep-seated
radioactivity!



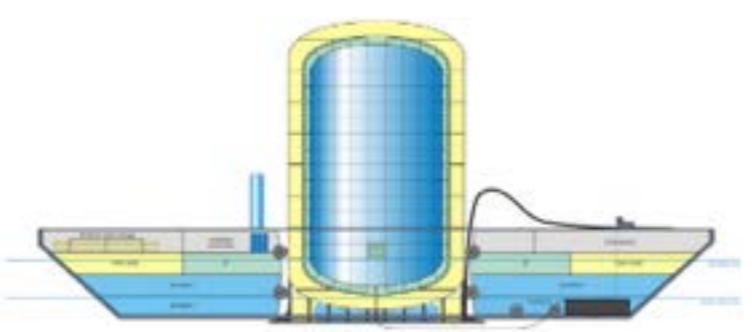
KamLAND (Japan)



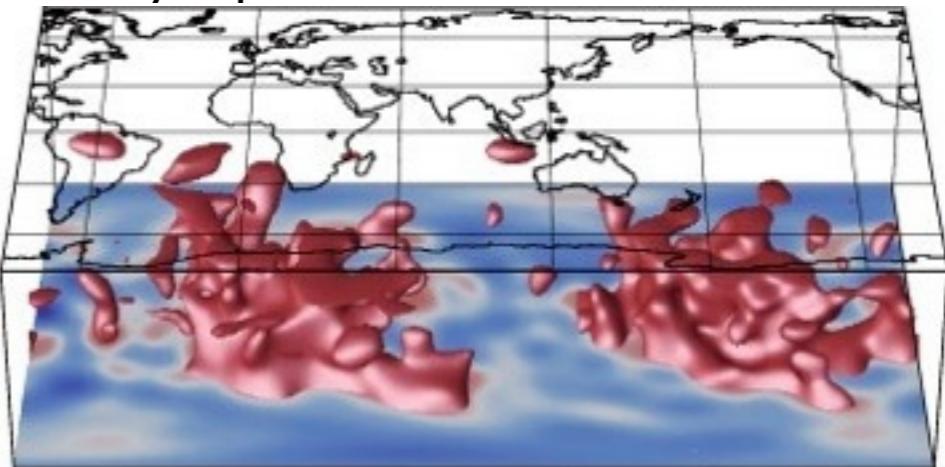
Borexino (Italy)



Hanohano, proposed
ocean deployment

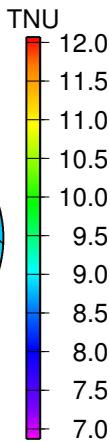
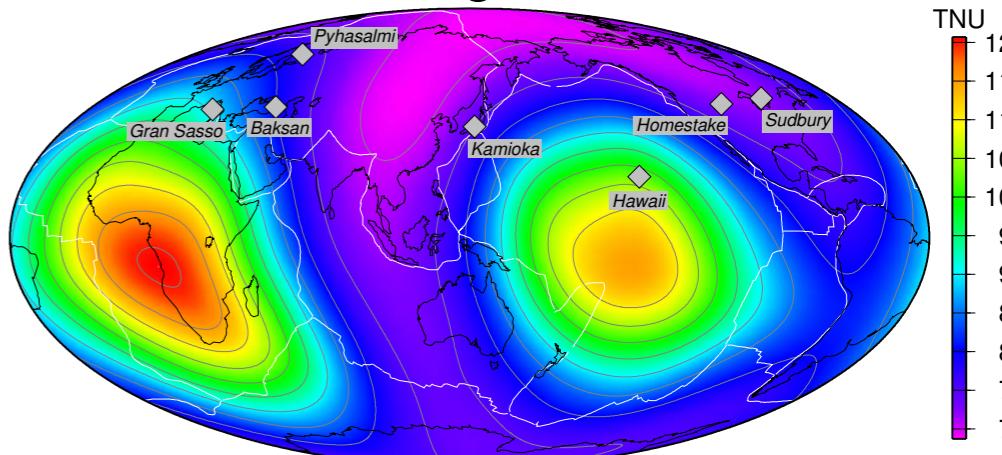


LLSVPs may represent material enriched in U,Th,K



from Bull et al. 2009

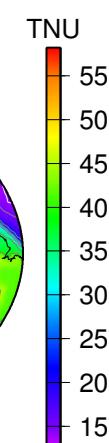
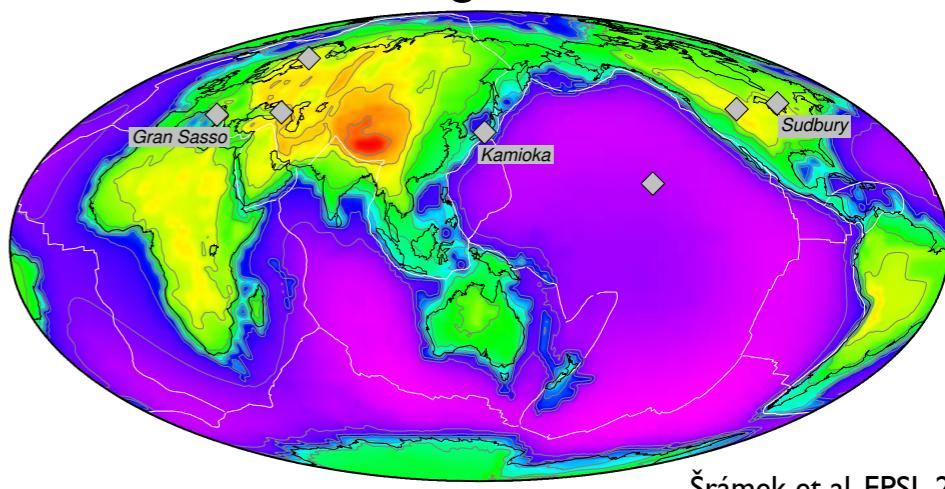
Predicted mantle geoneutrino flux



Šrámek et al. EPSL 2012

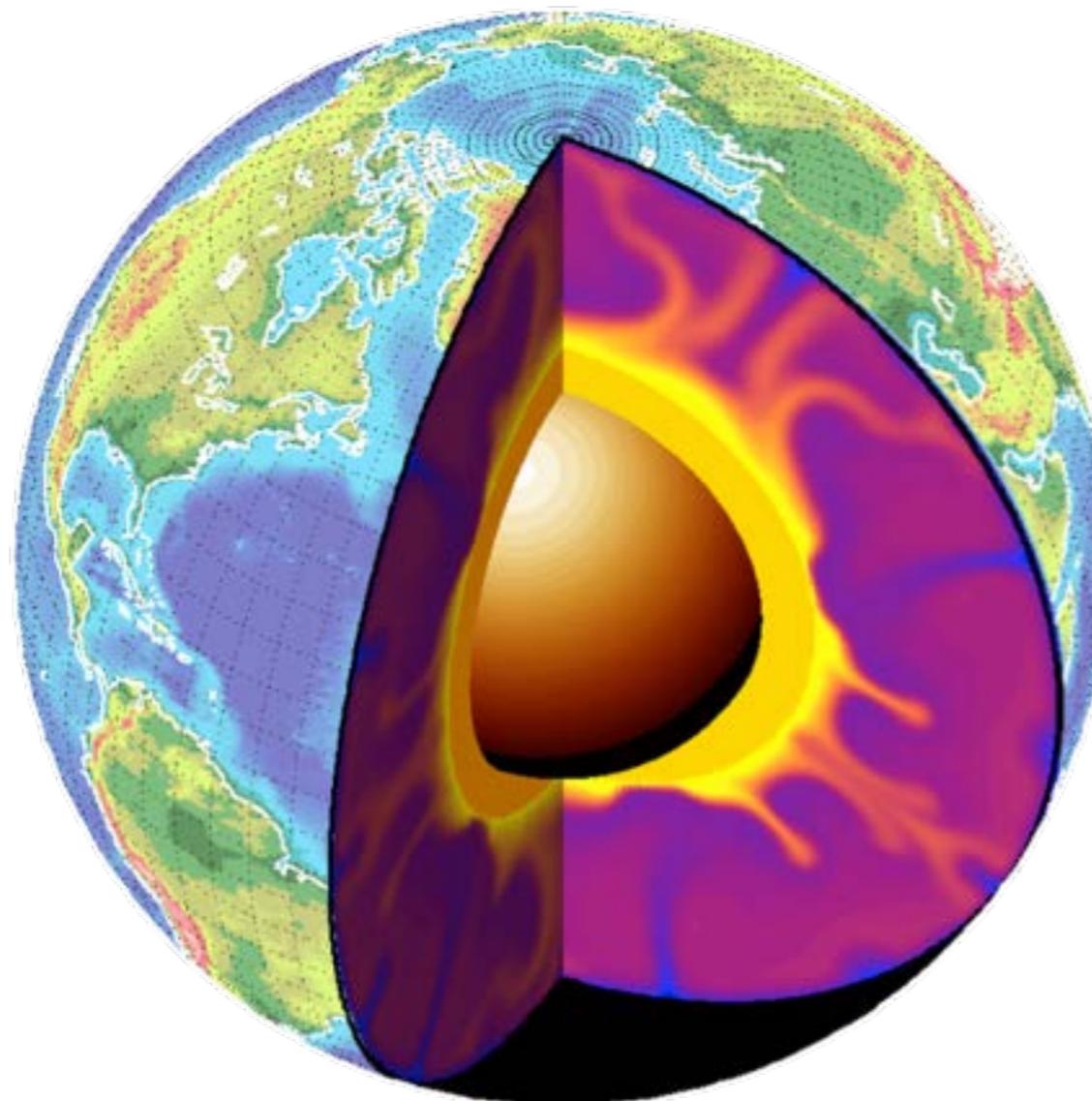
Mantle: detectable variation!
Must measure in the ocean...

Crust + mantle geoneutrino flux



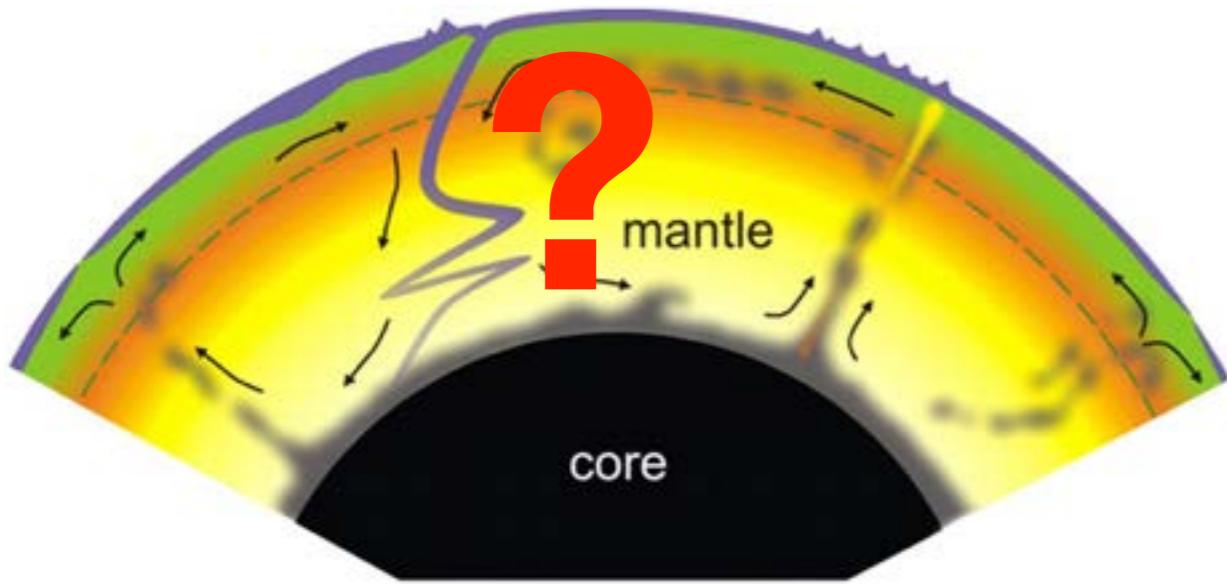
Šrámek et al. EPSL 2012

How much U, Th, K in the Earth?

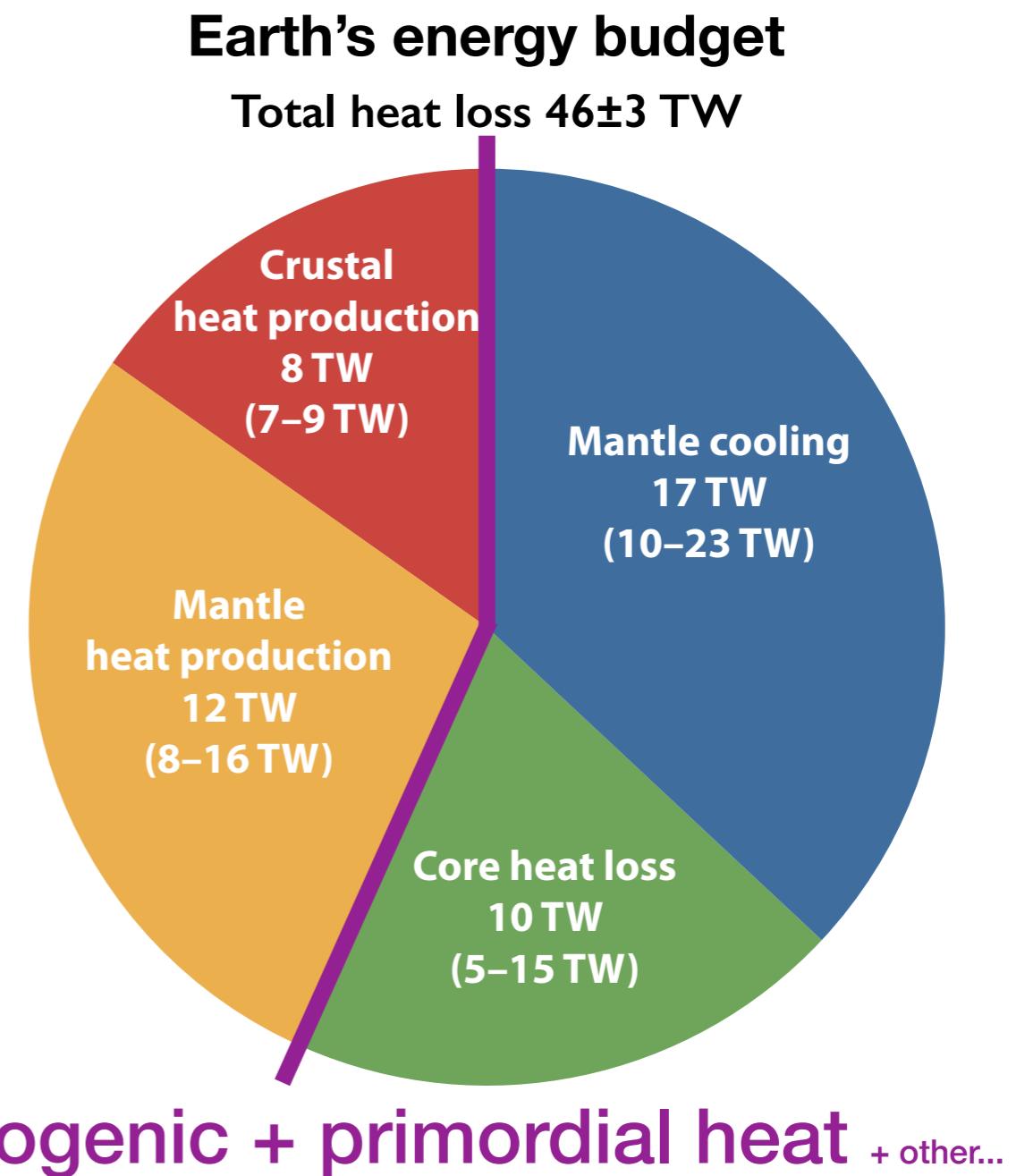


Estimates range from 9 to 36 TW radiogenic power

How much radiogenic heating in the mantle to power convection?



Estimates from 2 to 29 TW radiogenic power in the mantle



U Th K

Composition of Silicate Earth (BSE)

- “**Geochemical**” estimate
 - Ratios of RLE abundances constrained by C1 chondrites
 - Absolute abundances inferred from Earth rock samples
 - *McDonough & Sun (1995), Allègre (1995), Hart & Zindler (1986), Palme & O’Neill (2003), Arevalo et al. (2009)*
- “**Cosmochemical**” estimate
 - Isotopic similarity between Earth rocks and E-chondrides
 - Build the Earth from E-chondrite material
 - *Javoy et al. (2010)*
 - also “collisional erosion” models (*O’Neill & Palme 2008*)
- “**Geodynamical**” estimate
 - Based on a classical parameterized convection model
 - Requires a high mantle Urey ratio, i.e., high U, Th, K

TW radiogenic power

BSE

Mantle

20±4

13±4

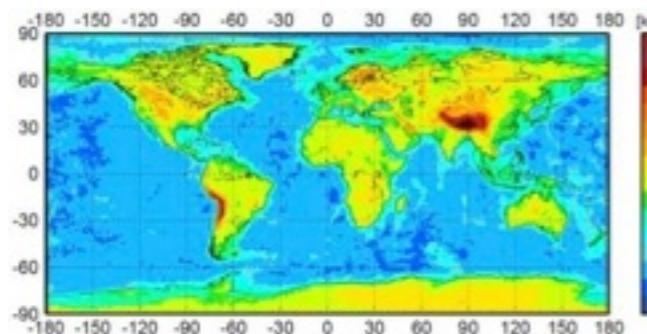
11±2

4±2

33±3

26±3

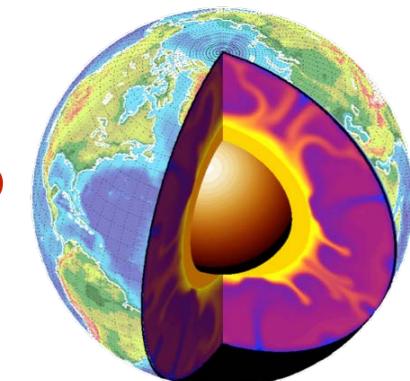
BSE = Mantle + Crust



Oceanic: 0.22 ± 0.03 TW

Continental: $6.8 (+1.4/-1.1)$ TW

?



Approach:

Use noble gas isotopic ratios
from gases originating at depth

- Why now?
- How to calculate ^{39}Ar production rate
- Some results

Isotopes of Argon

34 known isotopes

^{40}Ar ... radiogenic, stable

^{39}Ar ... cosmo/nucleogenic, $t_{1/2} = 269$ y

^{36}Ar ... primordial, stable

Atmosphere

^{40}Ar from degassing of Earth over 4.5 Gy

^{39}Ar produced cosmogenically from ^{40}Ar

$^{40}\text{Ar}/^{36}\text{Ar} = 295$

$^{39}\text{Ar}/^{40}\text{Ar} = 8 \times 10^{-16}$

Underground

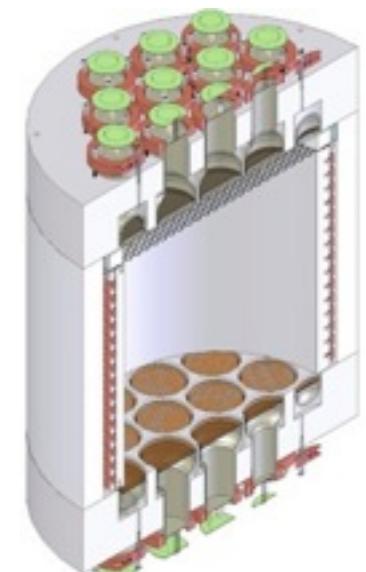
^{40}Ar produced by electron capture on ^{40}K

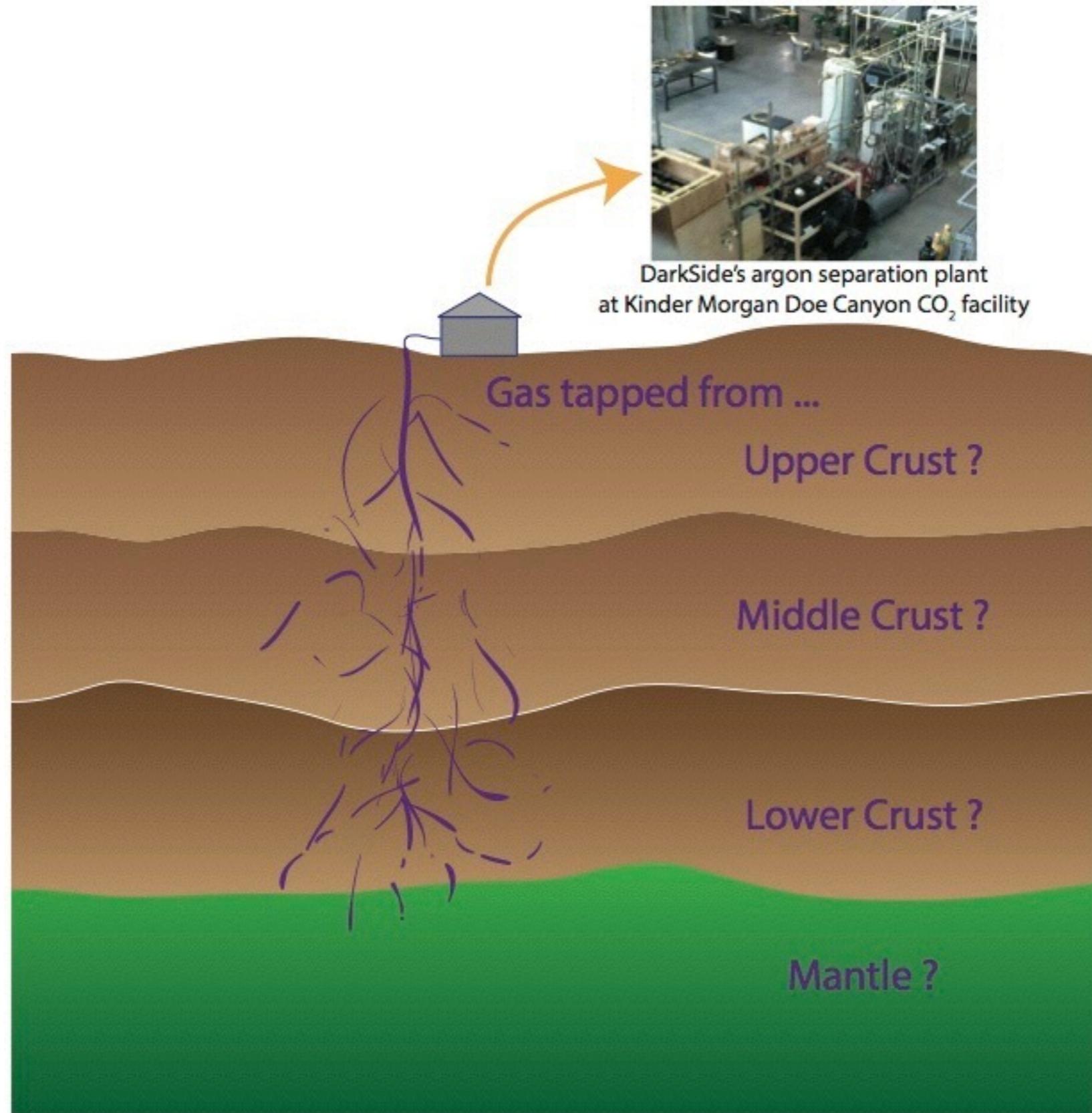
^{39}Ar produced nucleogenically from ^{39}K

$^{39}\text{Ar}/^{40}\text{Ar} < 0.006$ atmospheric

Dark matter WIMP search

- Dark matter detectors looking for Weakly Interacting Massive Particles (WIMPs) require low radioactivity argon
- Atmospheric level ($^{39}\text{Ar}/^{40}\text{Ar}=8\times 10^{-16}$) is too high
- Gas from deep CO₂ wells shows lower level of ^{39}Ar (e.g., Cortez CO, Bueyeros NM)
- $^{39}\text{Ar}/^{40}\text{Ar}$ – challenging measurements...
 - low-level radioactive decay counting
 - Atomic Trap Trace Analysis (ATTA)





Nuclear physics notation:

$A(a,b)B$

$A + a \rightarrow b + B$

A ... target nuclide

a ... projectile

b ... ejectile

B ... product nuclide

(a,n)

(n,p)

(n,a)

(n,2n)

...

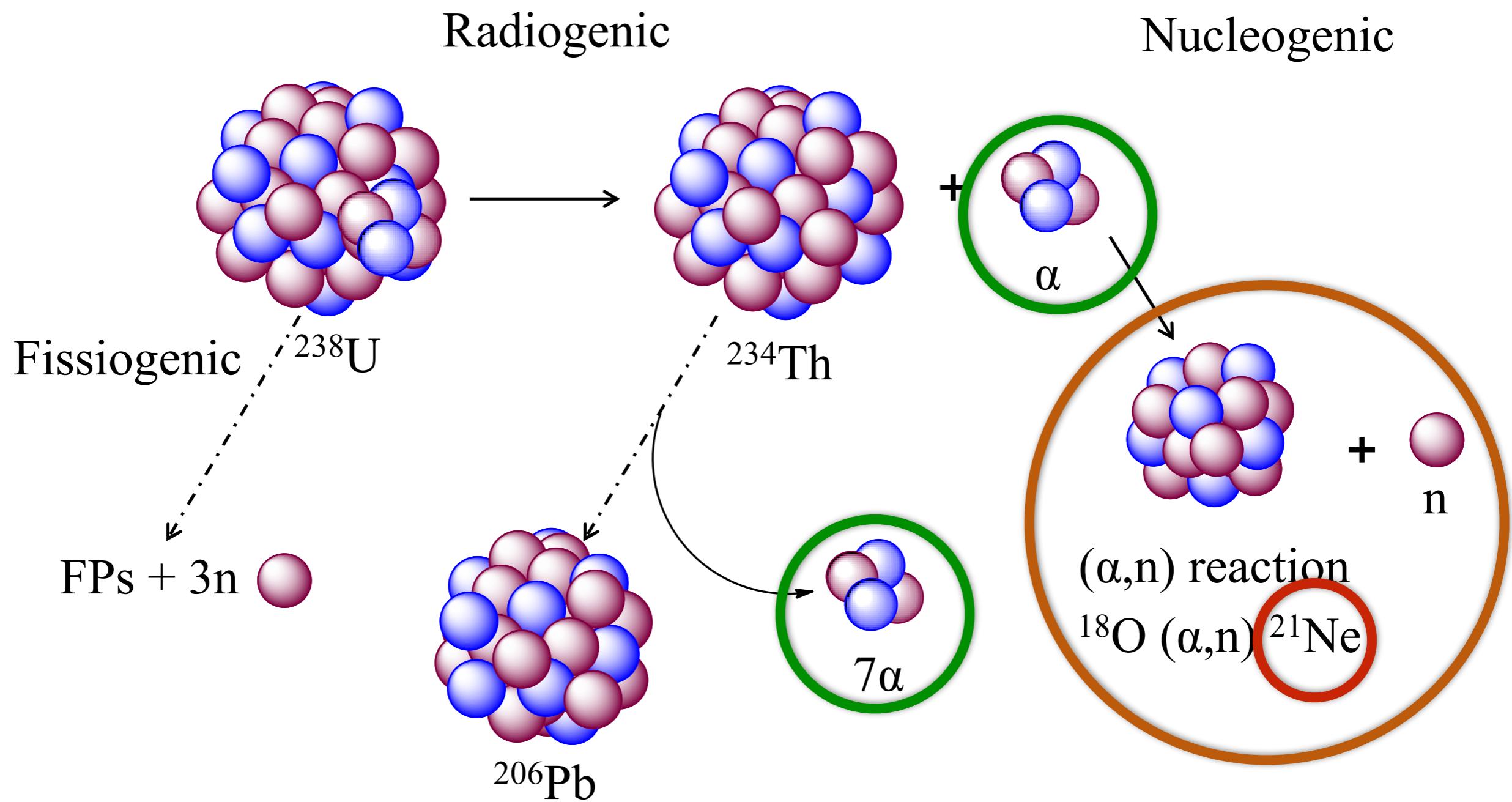
$^{40}\text{Ar}(n,2n)^{39}\text{Ar}$

atmosphere, cosmogenic

Nucleogenic production

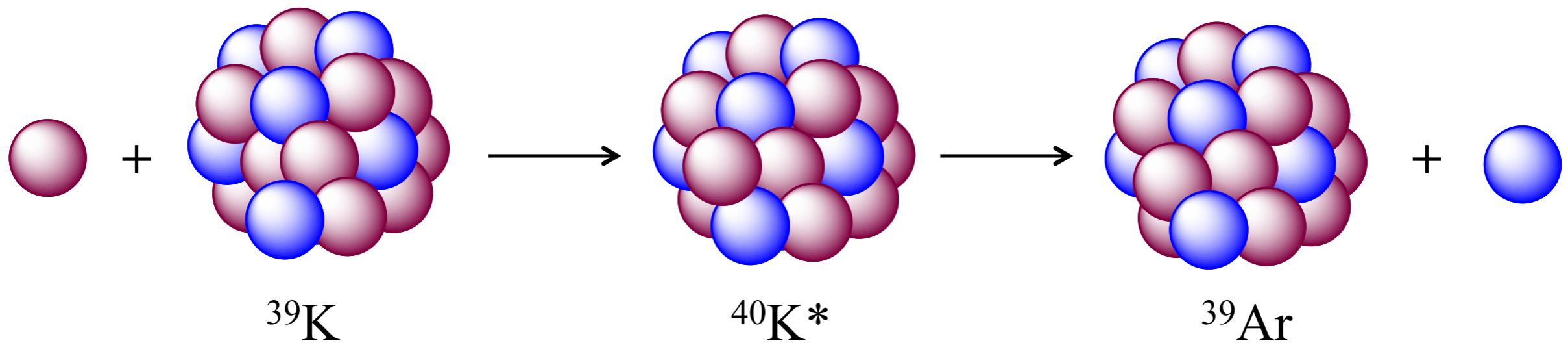
- Decay of radioactive U and Th in Earth's interior produce α particles
- (α, n) reactions on light isotopes produce neutrons
- Neutrons are also produced by spontaneous fission of U
- (n, p) reaction then produces ^{39}Ar from ^{39}K
- ⇒ Measurement of isotopic ratios in outgassing rock can inform us about the U, Th, K

1. U, Th decay produces α's



2. (α,n) produce neutrons

$^{39}\text{K}(\text{n},\text{p})$ produces ^{39}Ar



composite
nucleus

Noble gas isotopic ratios

- ^{21}Ne production rate proportional to [U+Th]
- ^{39}Ar production rate proportional to [K] \times [U+Th]
- ^{40}Ar production rate proportional to [K]

therefore

- $^{39}\text{Ar}/^{40}\text{Ar}$ proportional to [U+Th]
- $^{39}\text{Ar}/^{21}\text{Ne}$ proportional to [K]
- $^{40}\text{Ar}/^{21}\text{Ne}$ proportional to [K]/[U+Th]

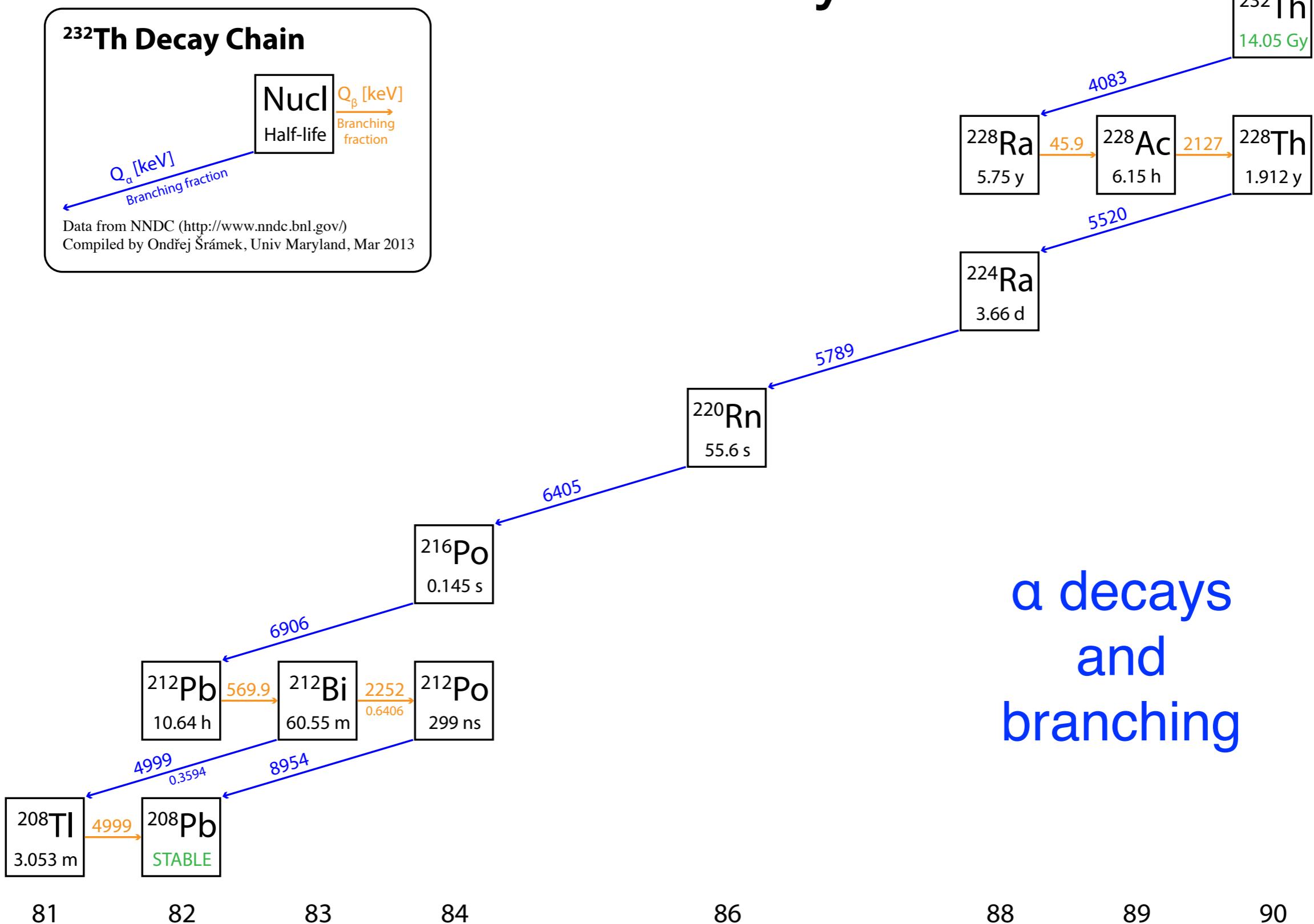
Calculating ^{39}Ar nucleogenic production rate

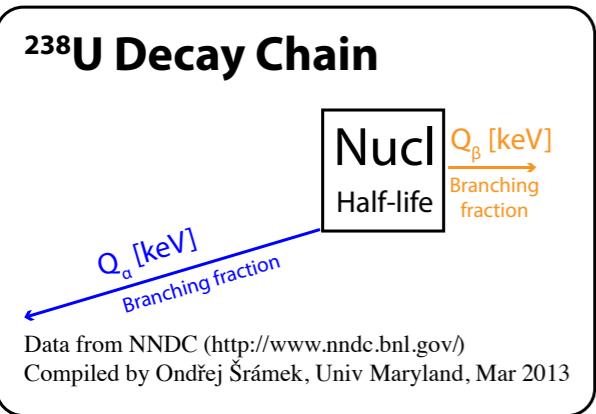
- α emission from natural radionuclides
- neutron production by α -induced reactions
- $^{39}\text{K}(\text{n},\text{p})^{39}\text{Ar}$

Table 7: ^{39}Ar production rates as calculated in several studies. Rates are recalculated to a common K, Th, U composition of Upper Crust in *Rudnick and Gao (2003)*, K=2.3 %, Th=10.5 ppm, U=2.7 ppm by weight. ^{39}Ar prod. rate in number of atoms per year per kg of rock.

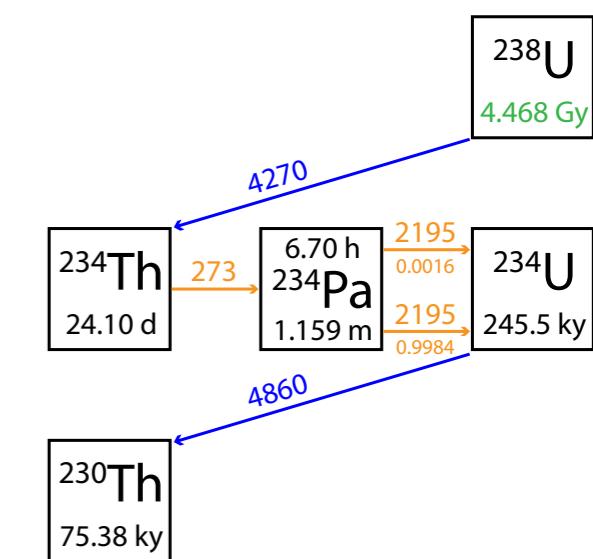
| Reference | ^{39}Ar prod. rate |
|------------------------------|-----------------------------|
| <i>Mei et al. (2010)</i> | 11 |
| <i>Yokochi et al. (2012)</i> | 55 |
| <i>Yokochi et al. (2013)</i> | 170; 110 |
| This study | 30 |

^{232}Th decay chain





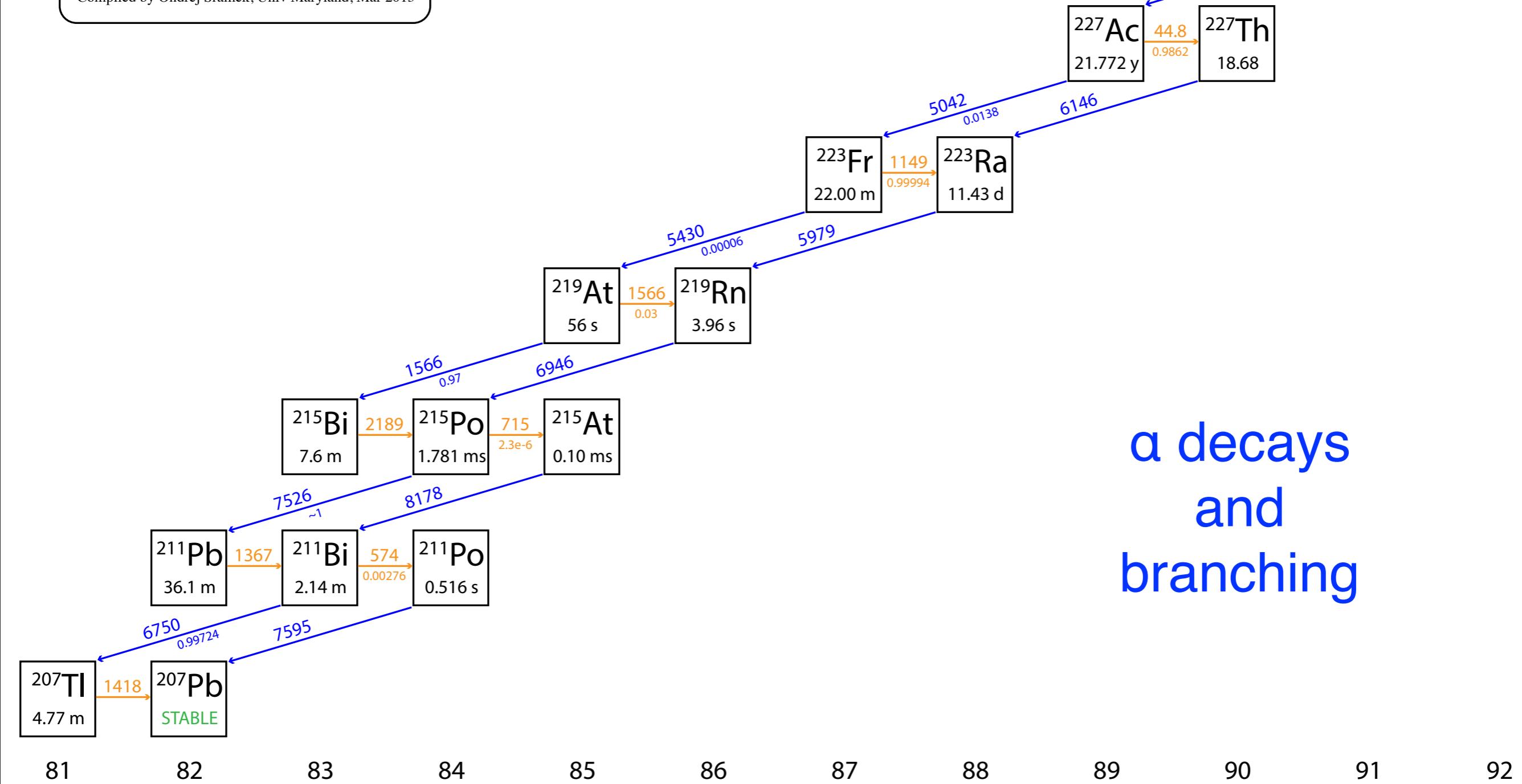
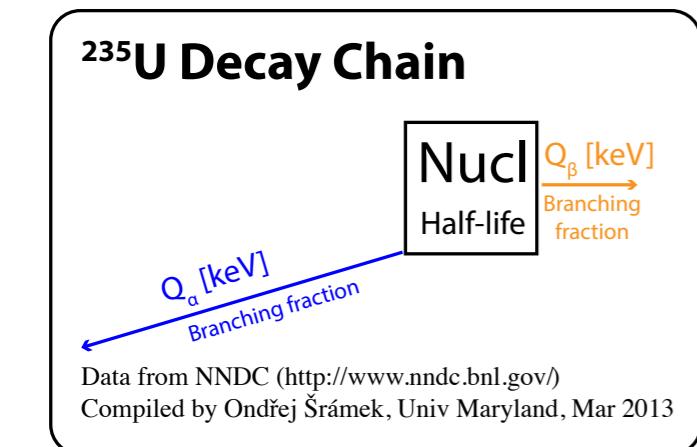
^{238}U decay chain



a decays
and
branching

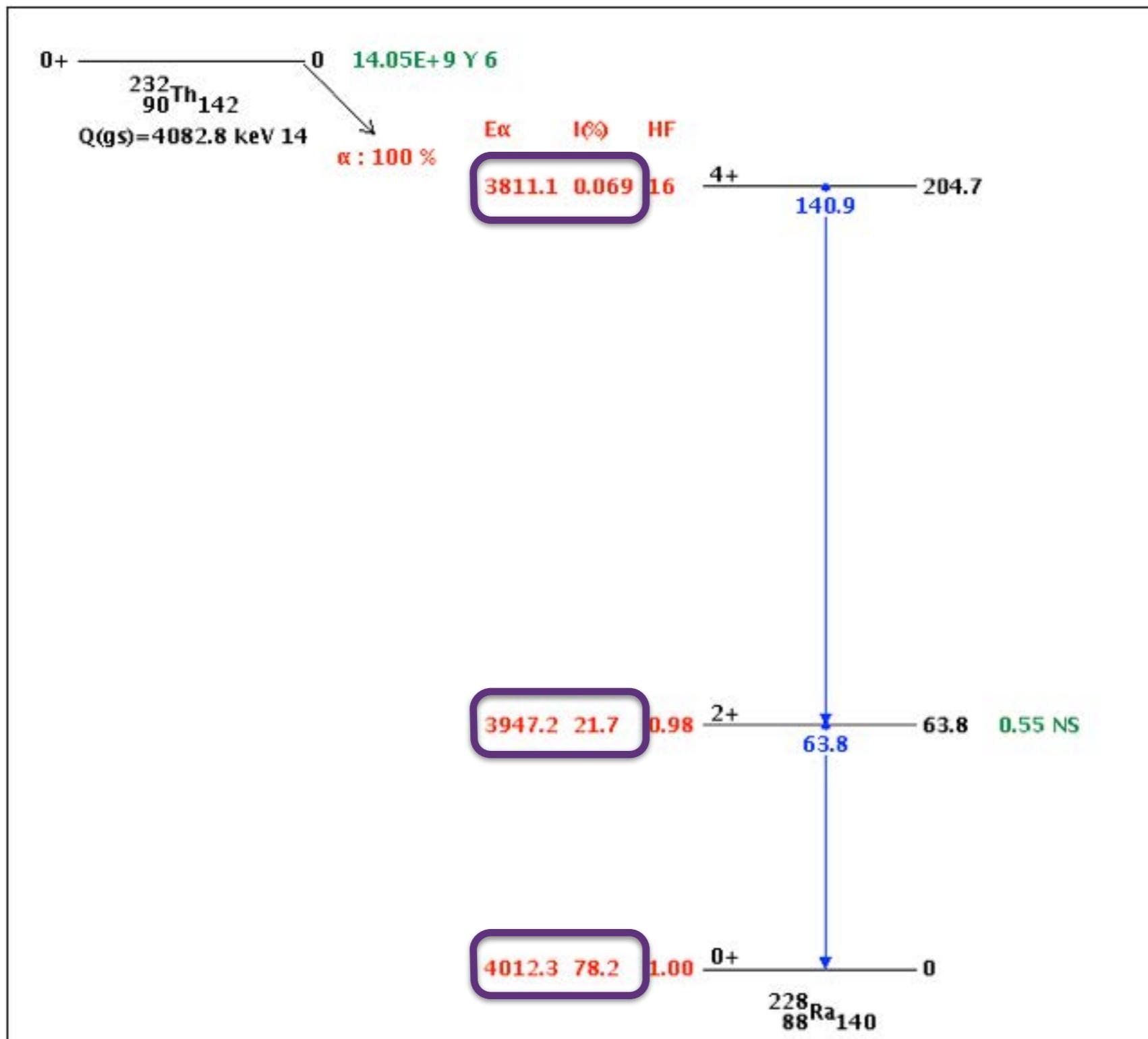
^{235}U
703.8 My

^{235}U decay chain



a decays
and
branching

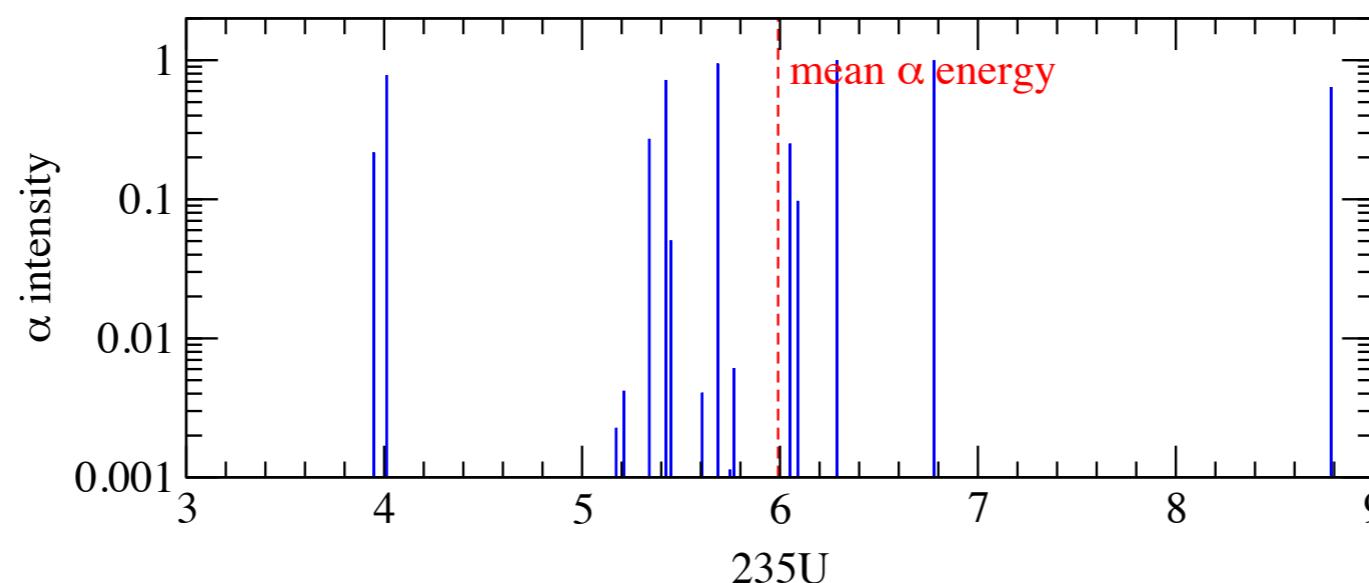
^{232}Th α decay scheme



α energies
and
intensities

Natural alpha emission energy spectra

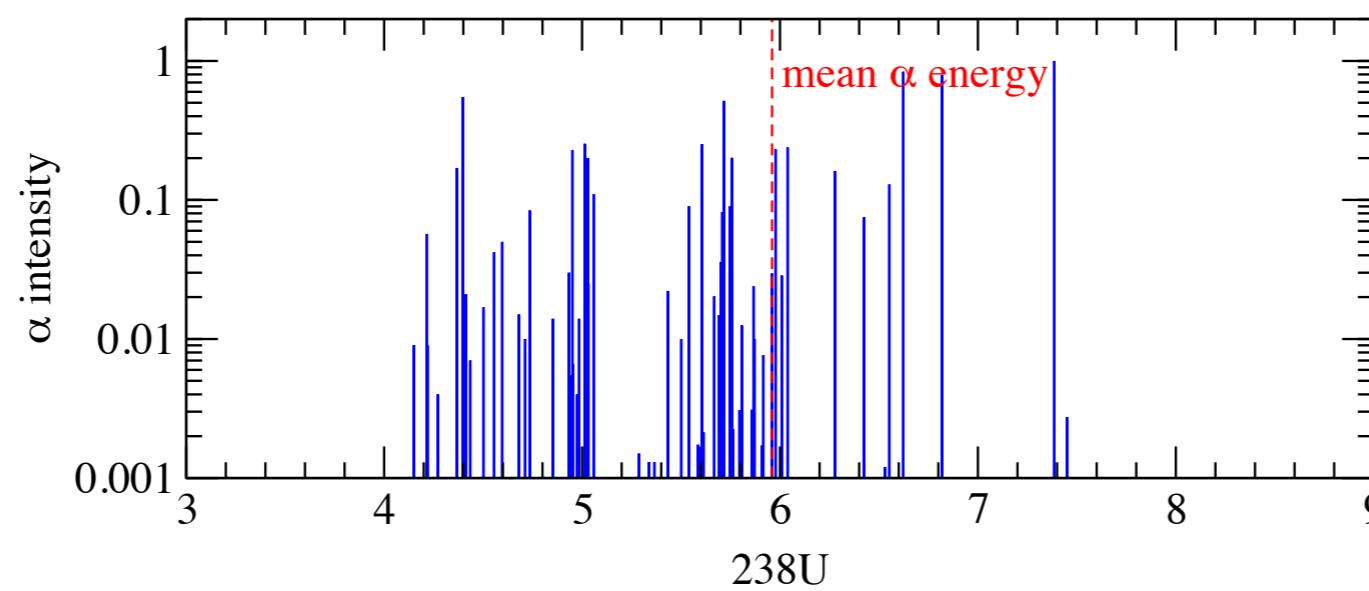
232Th



a's/decay:

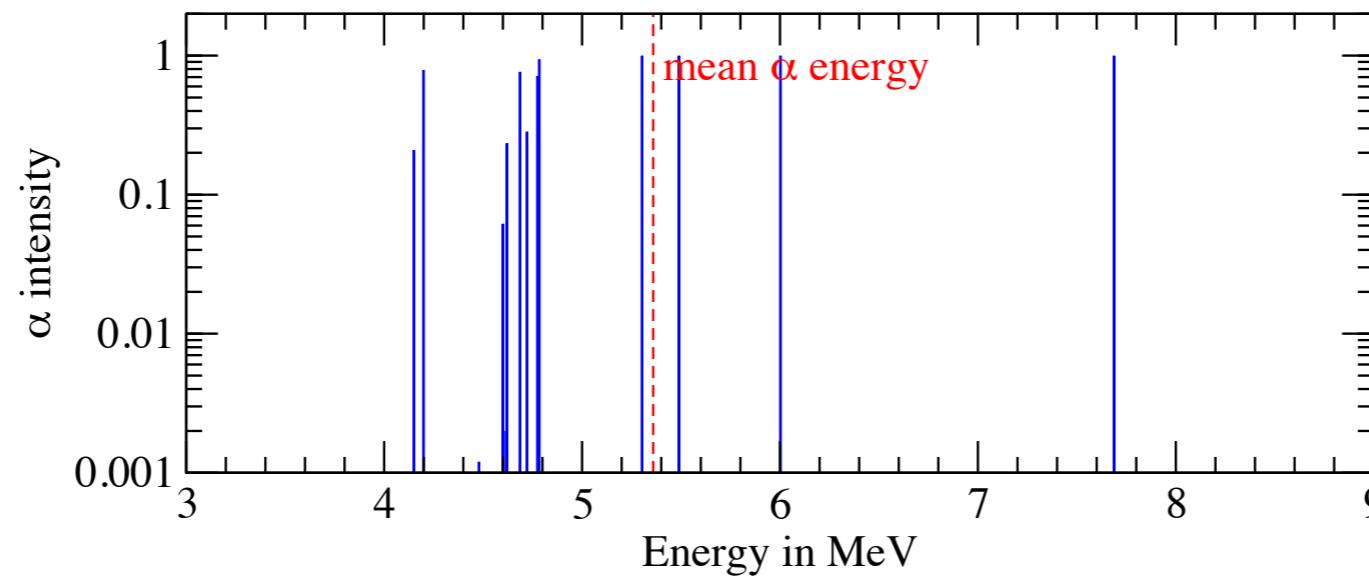
6

235U



7

238U



8

Energy in MeV

Calculating ^{39}Ar nucleogenic production rate

- α emission from natural radionuclides
- neutron production by α -induced reactions
- $^{39}\text{K}(\text{n},\text{p})^{39}\text{Ar}$

(a,n) neutron production and spectrum

- α emitted at initial energy $E_{\alpha 0}$
- α slows down and eventually stops ($E=0$) in the rock
- before it stops, it can participate in (a,n) reaction

neutron production function
neutrons per 1 α particle

$$P_i(E_{\alpha 0}) = N_i \int_0^{E_{\alpha 0}} \frac{\sigma_{\alpha, n}^i(E_\alpha)}{\left(-\frac{dE_\alpha}{dx}\right)} dE_\alpha$$

↑
atomic density ↑
 cross section
 ↑
 stopping power

neutron spectrum

$$\frac{dP_i}{dE_n}(E_{\alpha 0}, E_n) = N_i \int_0^{E_{\alpha 0}} \frac{\frac{d\sigma_{\alpha, n}^i}{dE_n}(E_\alpha, E_n)}{\left(-\frac{dE_\alpha}{dx}\right)} dE_\alpha$$

↑
spectrum (or differential c.s.)

(a,n) neutron yield and production rate

neutron production function

neutrons per 1 a particle

neutron spectrum

neutron yield

neutrons per decay of 1 atom of parent nuclide

neutron production rate

neutrons per unit time per unit mass of rock

$$P_i(E_{\alpha_0}) = N_i \int_0^{E_{\alpha_0}} \frac{\sigma_{\alpha,n}^i(E_\alpha)}{(-\frac{dE_\alpha}{dx})} dE_\alpha$$

cross section

atomic density

stopping power

spectrum (or differential c.s.)



$$\frac{dP_i}{dE_n}(E_{\alpha_0}, E_n) = N_i \int_0^{E_{\alpha_0}} \frac{\frac{d\sigma_{\alpha,n}^i}{dE_n}(E_\alpha, E_n)}{\left(-\frac{dE_\alpha}{dx}\right)} dE_\alpha$$

$$Y_{(\alpha,n)} = \sum_{k=1}^K b_k \sum_{l=1}^{L_k} f_l^\alpha \sum_{i=1}^I P_i(E_l)$$

a decays E levels targets
 branching ratio level intensity

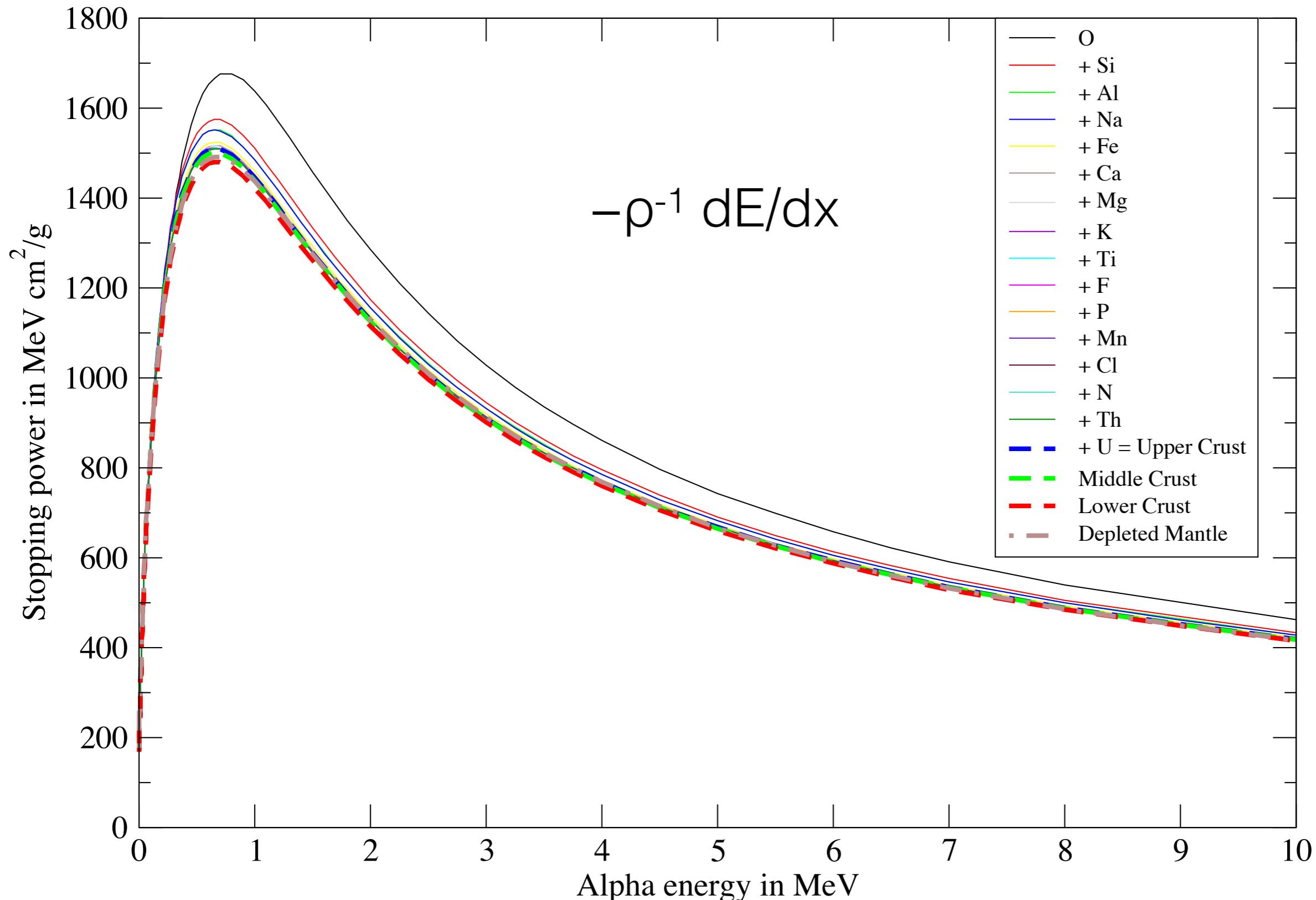
$$S_n = \lambda \frac{AXN_A}{M} Y$$

↓ ↓ ↓ ↓ ↓

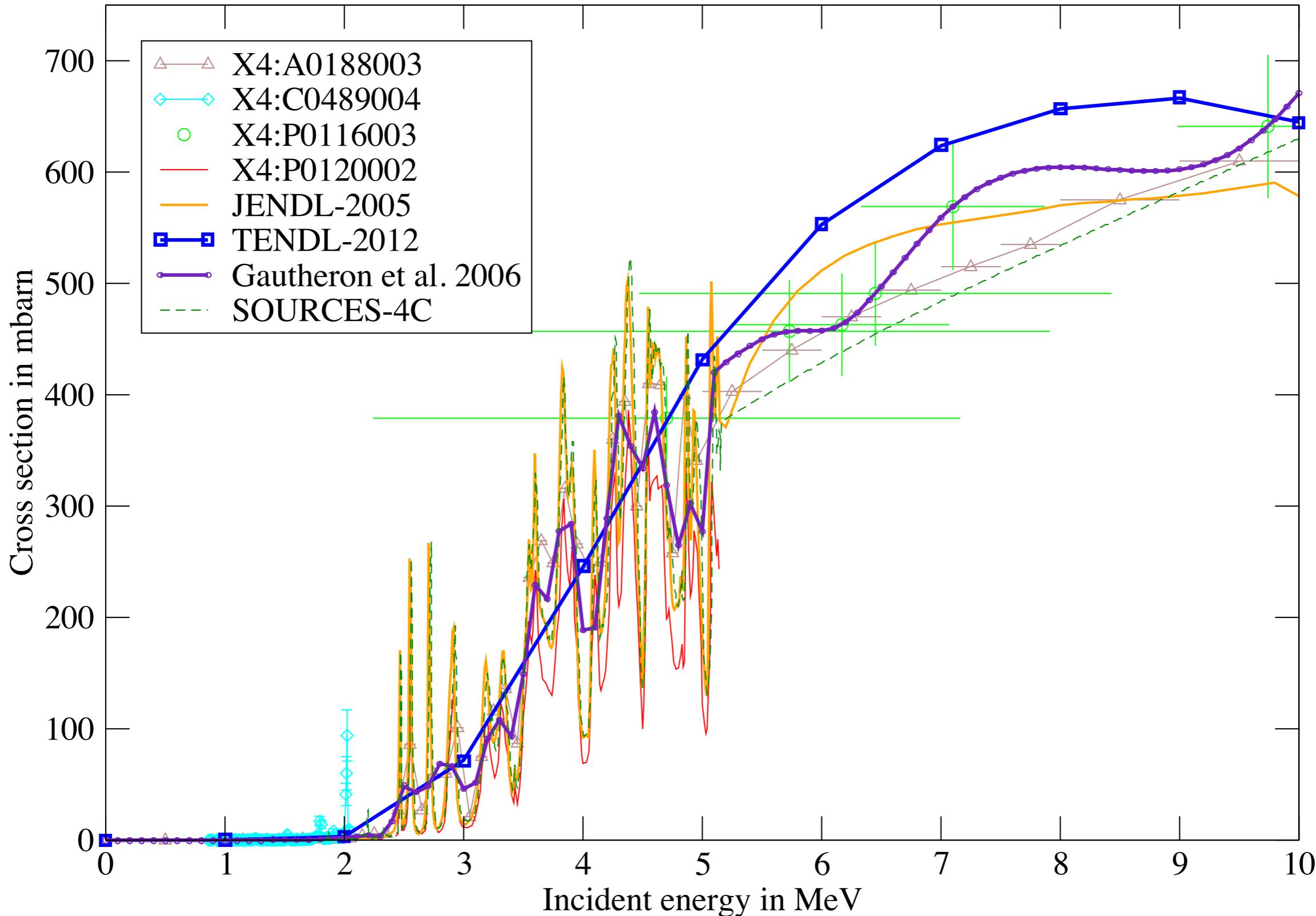
elem. abundance nat.isot.rat.
 decay constant atomic mass

Stopping power of alphas in rock

SRIM calc, RG03 / SS04 composition and CRUST2.0 / PREM density (2.70 - 2.88 - 3.05 / 3.42 g/cm³)

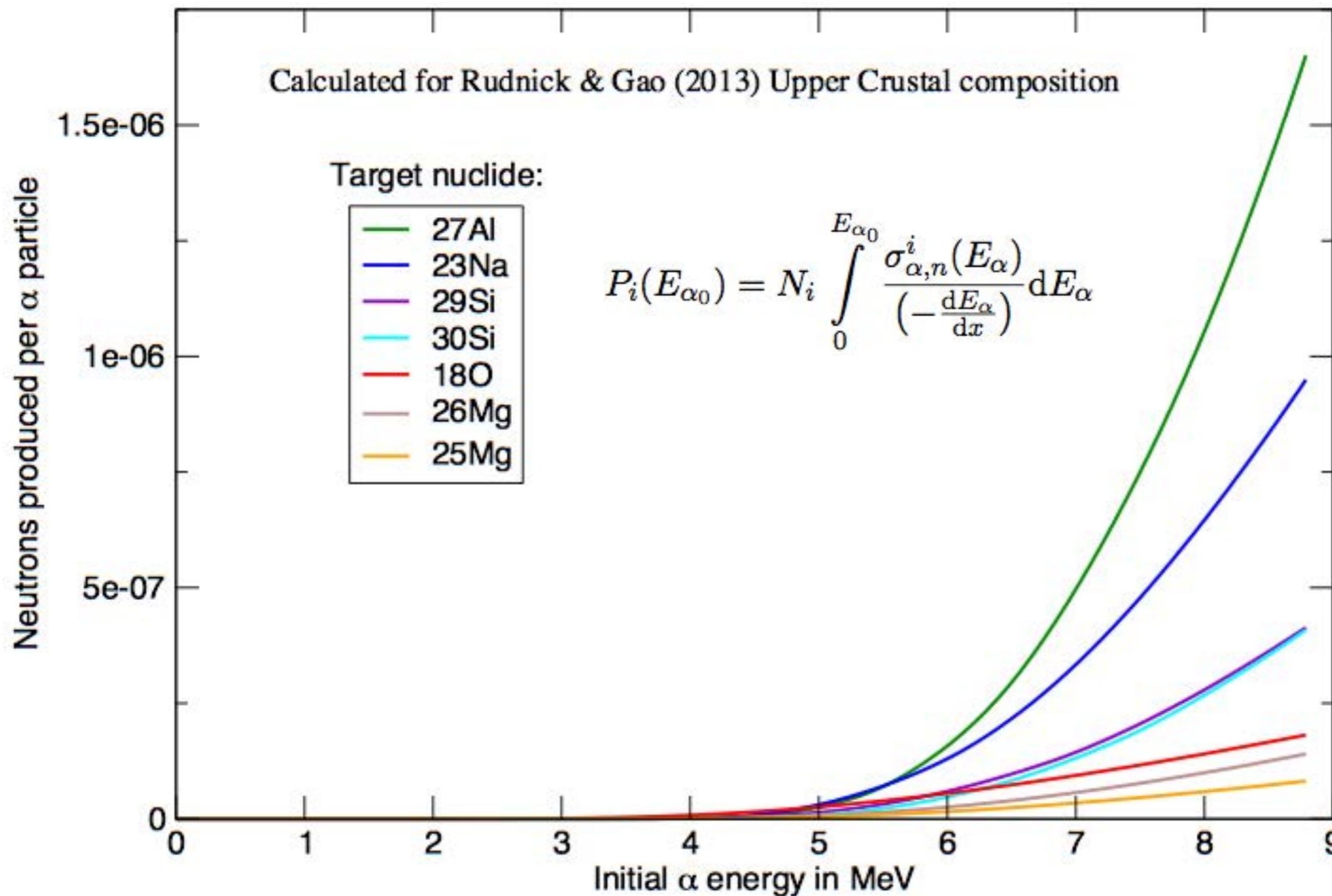


$^{18}\text{O}(\text{a},\text{n})$ cross section



We use: cross sections calculated by TALYS code, <http://www.talys.eu/>

Neutron production function



Calculating neutron spectra

1. Non-relativistic reaction kinematics Ground state-to-ground state

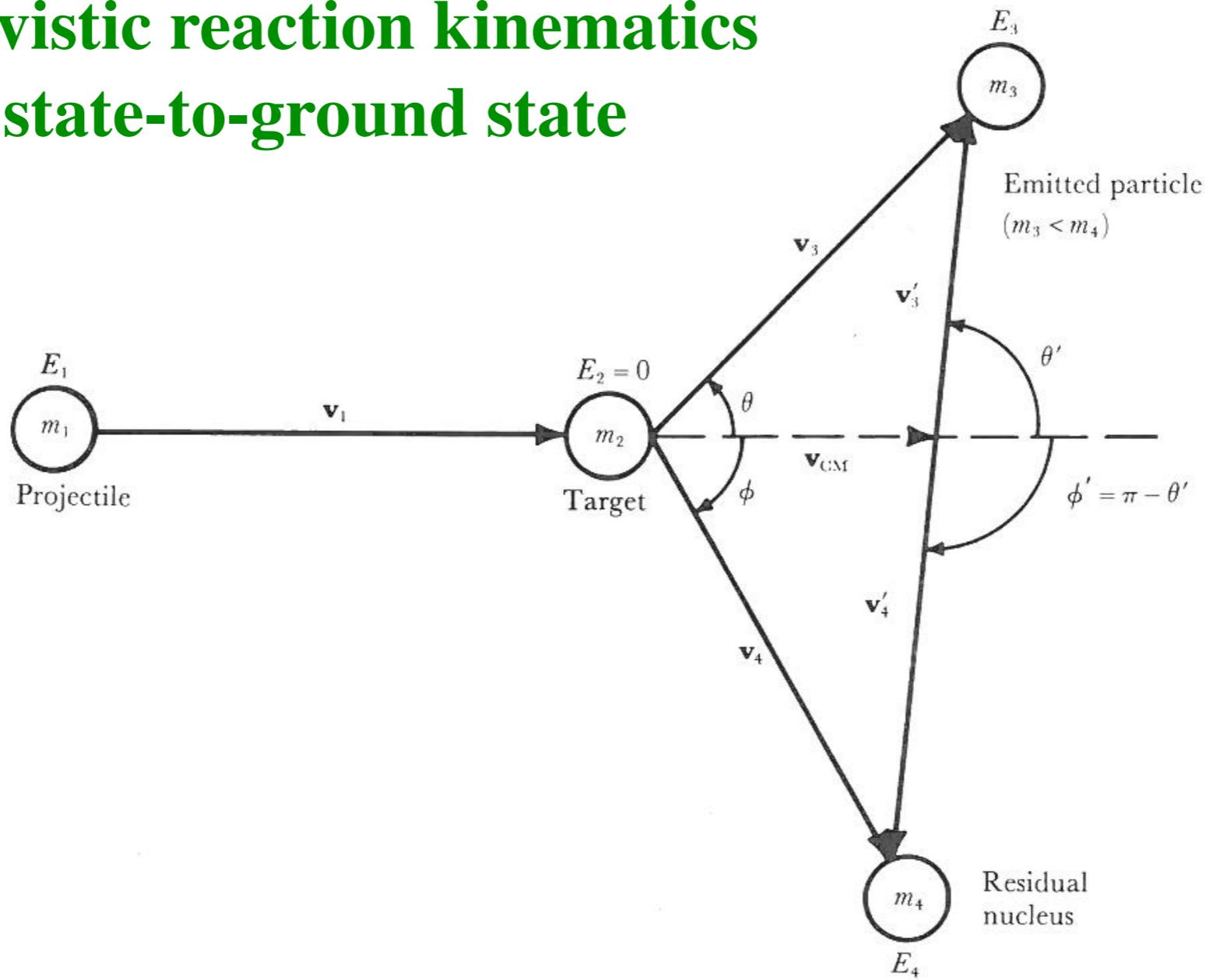
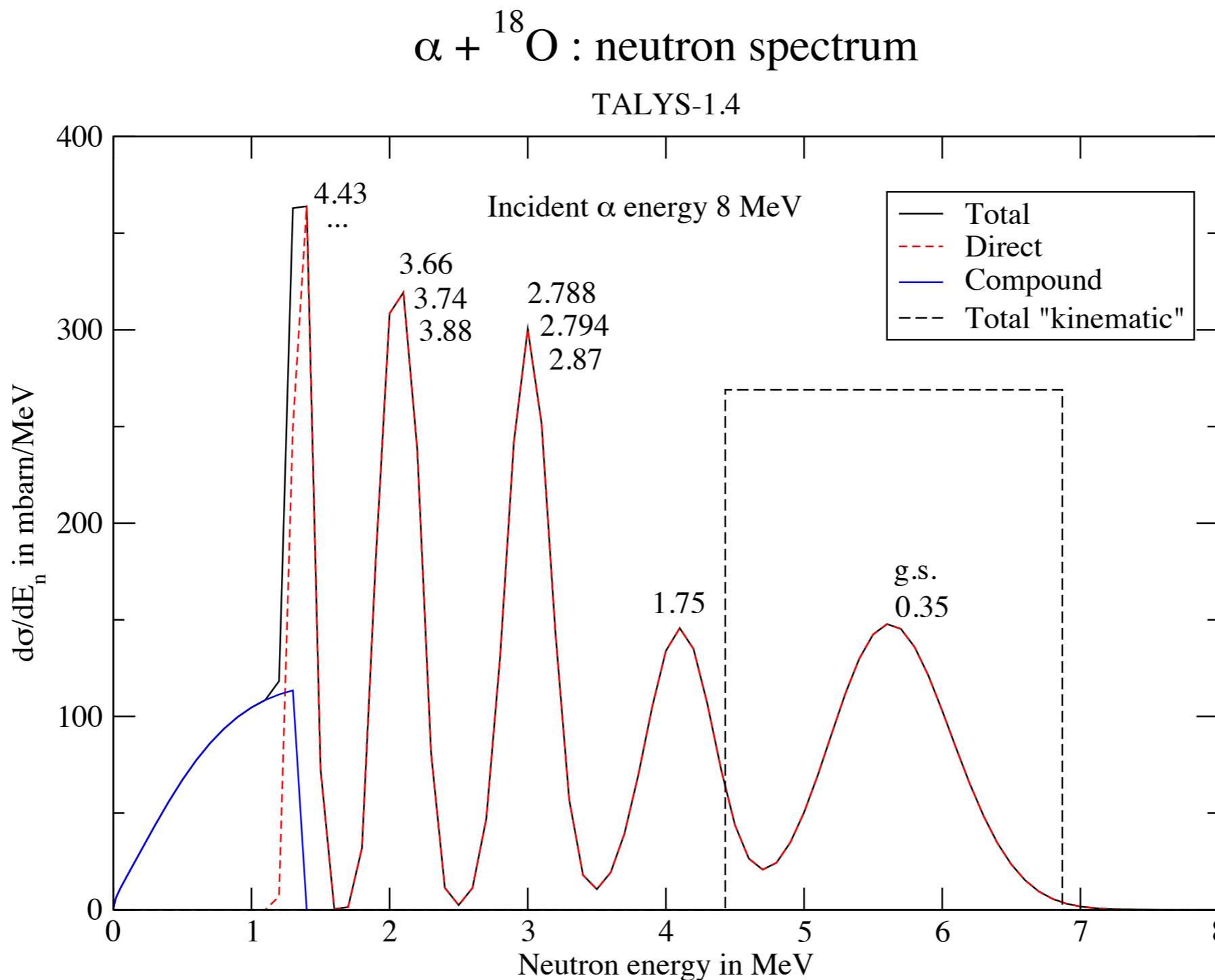


Fig. C-4. Vector diagram showing the relationship of kinematic quantities in the reaction process $m_1 + m_2 \rightarrow m_3 + m_4$.

2. More complete physical picture: TALYS code

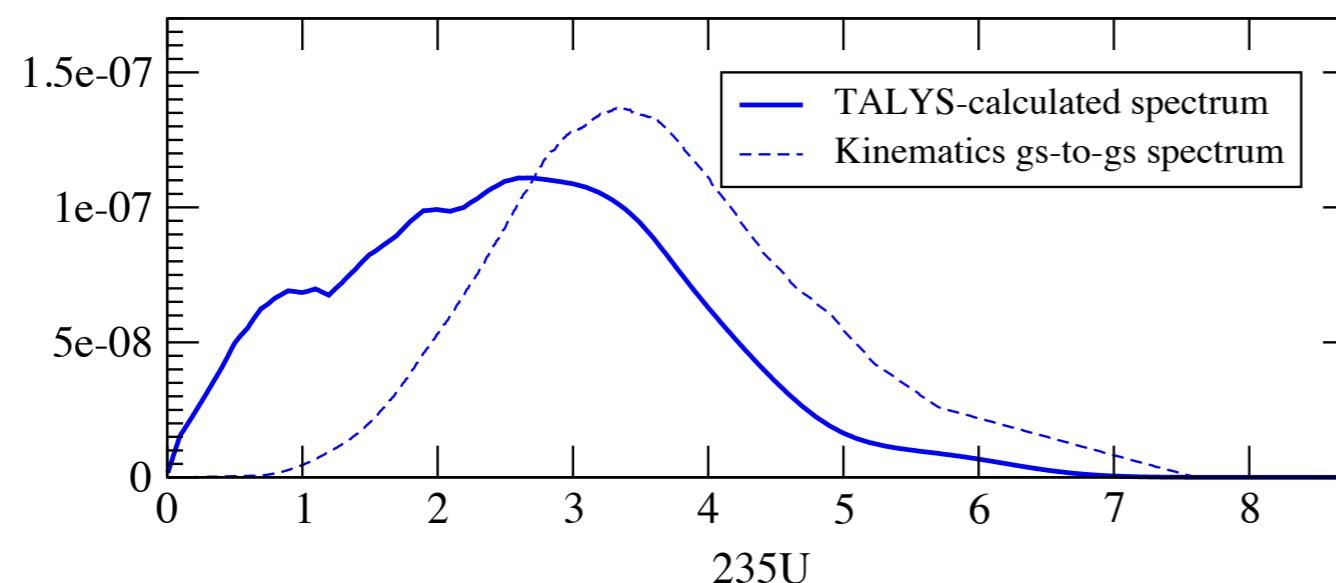
Neutron spectrum for 8 MeV α + ^{18}O



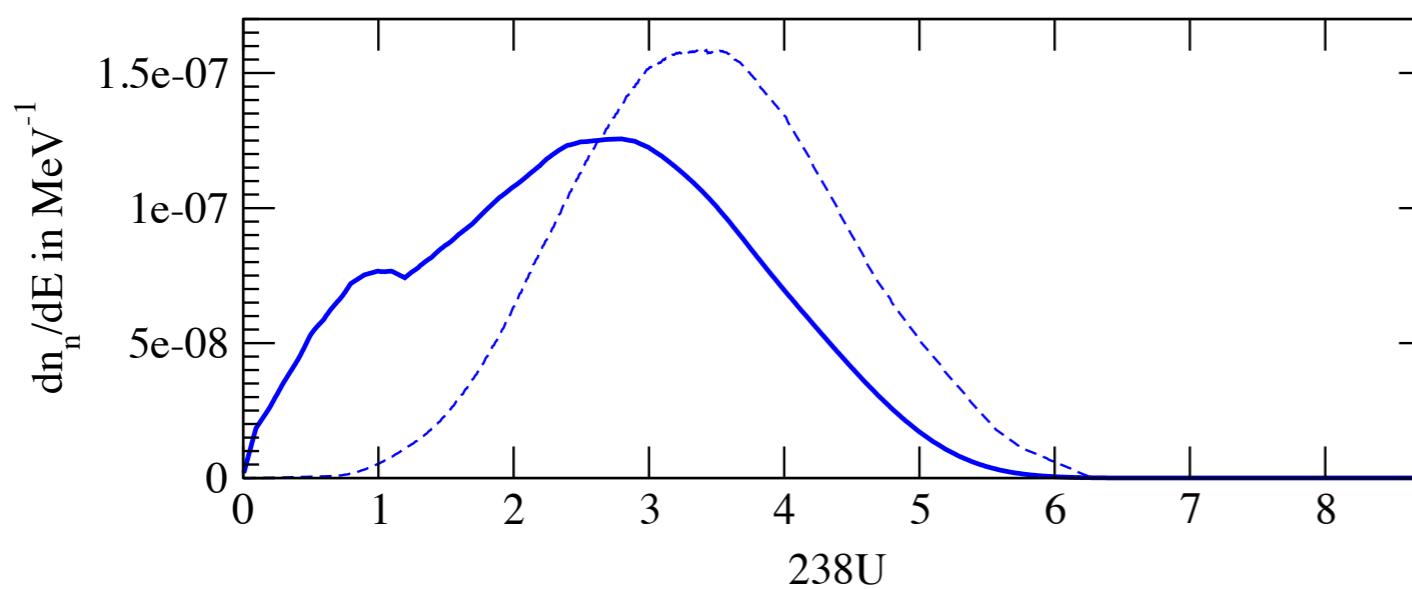
We use: spectra calculated by TALYS code, <http://www.talys.eu/>

$\alpha + {}^{18}\text{O}$: neutron energy spectra

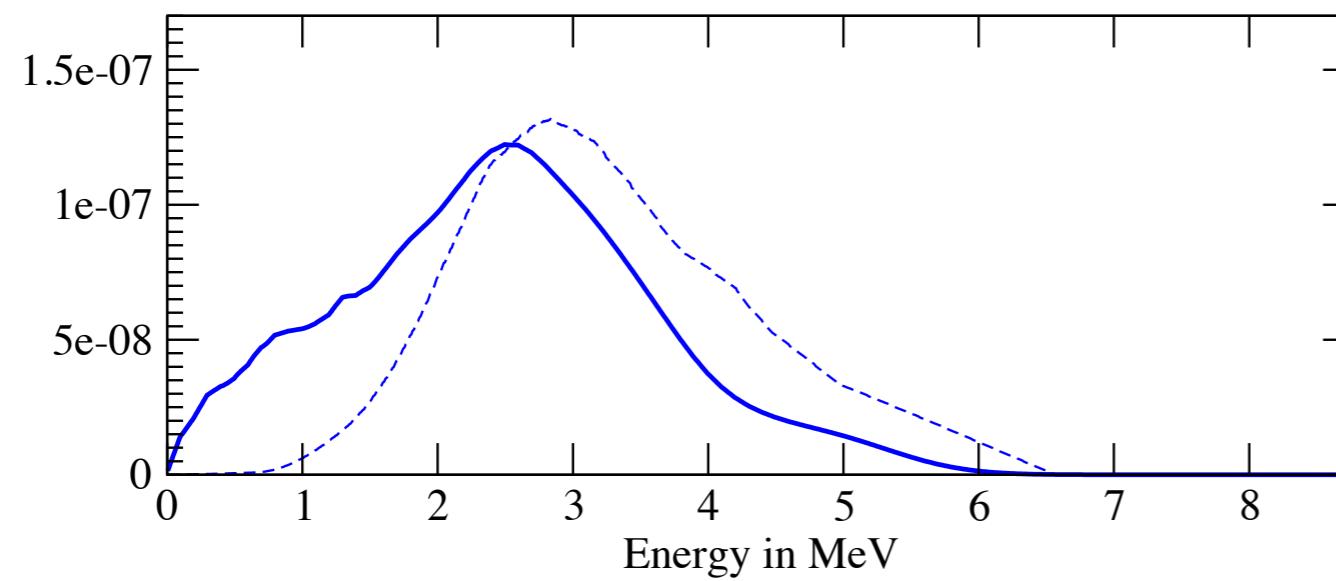
232Th



235U



238U

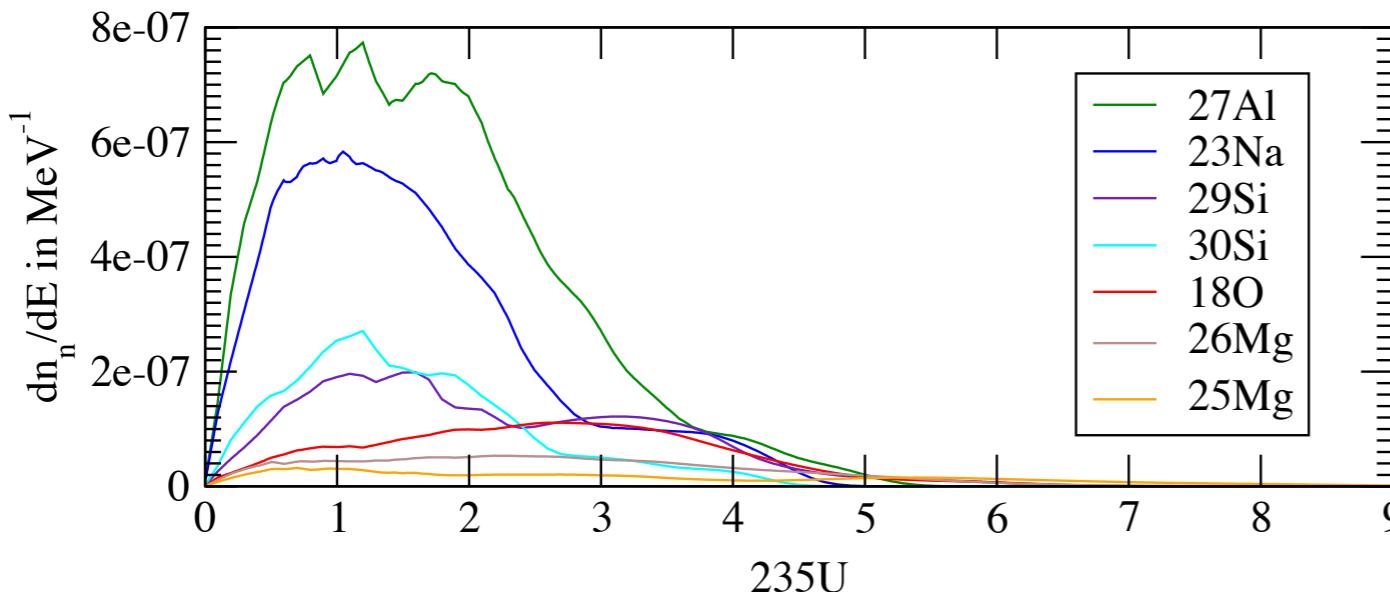


$$\frac{dP_i}{dE_n}(E_{\alpha_0}, E_n) = N_i \int_0^{E_{\alpha_0}} \frac{\frac{d\sigma_{\alpha,n}^i}{dE_n}(E_\alpha, E_n)}{\left(-\frac{dE_\alpha}{dx}\right)} dE_\alpha$$

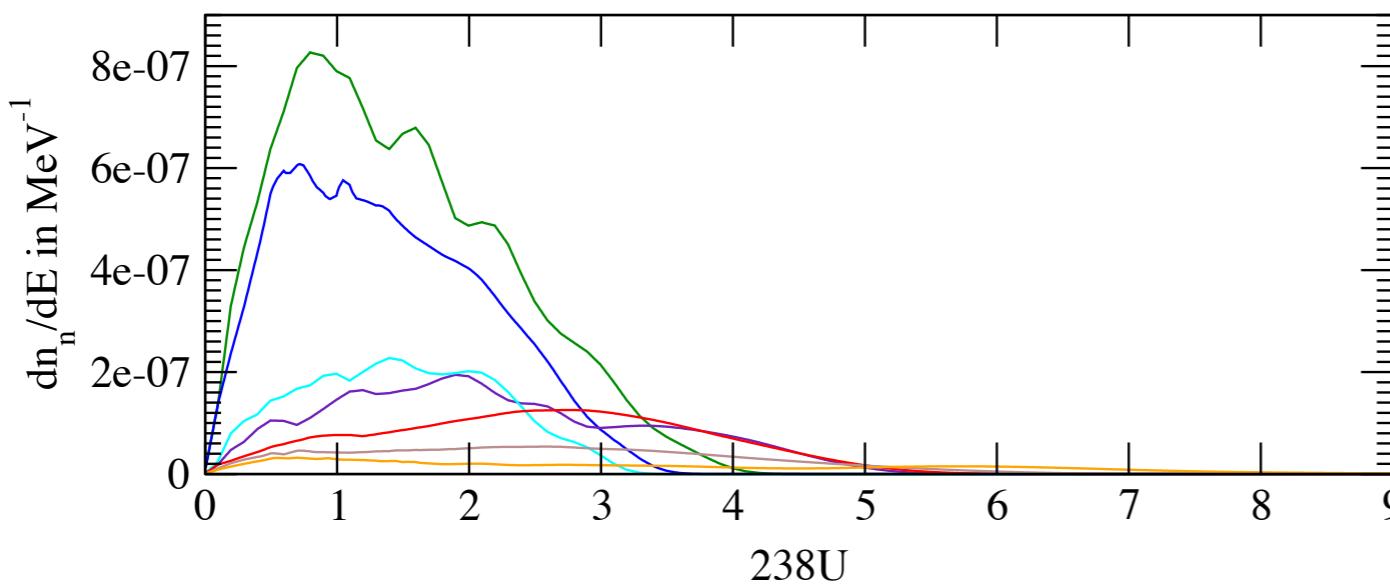
+ sum up over all α -decays,
all α levels

α + target : neutron energy spectra [TALYS]

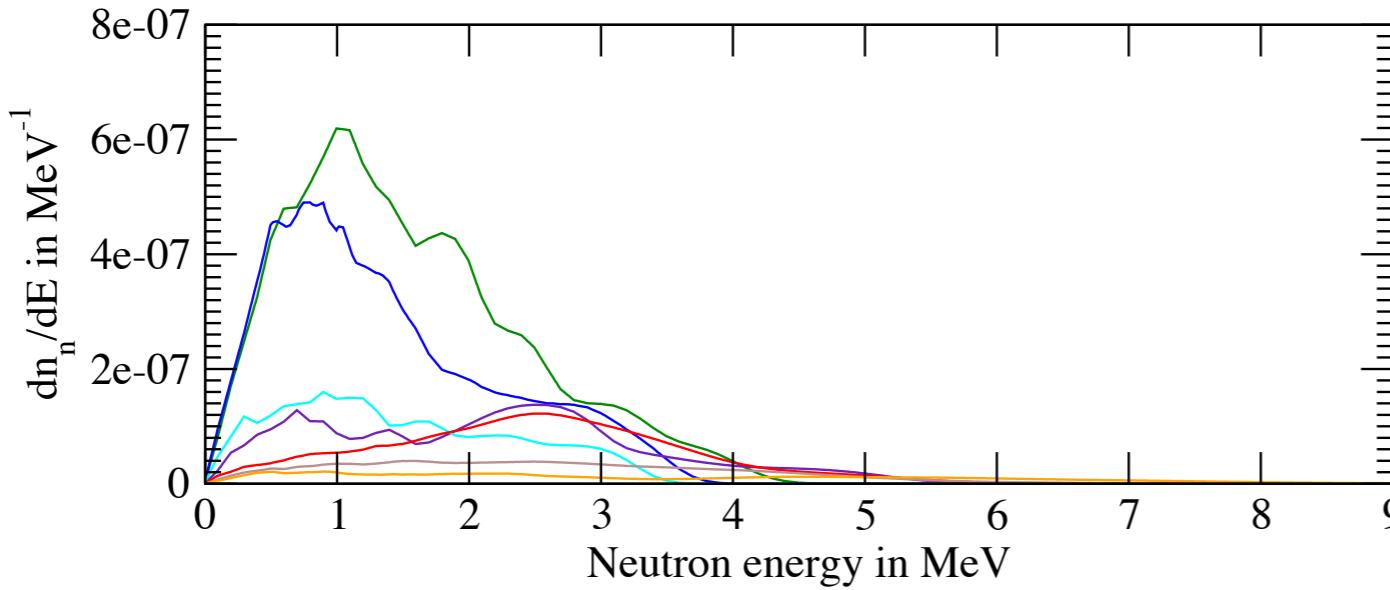
232Th



235U



238U



$$\frac{dP_i}{dE_n}(E_{\alpha_0}, E_n) = N_i \int_0^{E_{\alpha_0}} \frac{\frac{d\sigma_{\alpha,n}^i}{dE_n}(E_{\alpha}, E_n)}{\left(-\frac{dE_{\alpha}}{dx}\right)} dE_{\alpha}$$

+ sum up over all α -decays,
all α levels

Calculating ^{39}Ar nucleogenic production rate

- α emission from natural radionuclides
- neutron production by α -induced reactions
- $^{39}\text{K}(\text{n},\text{p})^{39}\text{Ar}$, also $^{24}\text{Mg}(\text{n},\alpha)^{21}\text{Ne}$

Neutrons propagating and interacting in the rock

- Competition between neutron scattering, various neutron-induced reactions on various targets
- We are interested in specific reactions
- Back-of-the-envelope calculations give us an order of magnitude answer:

$$P_{39}(E_0) = L_n(E_0) \bar{\sigma}_{39} N_{39} \quad (22)$$

where $L_n(E_0)$ is the distance traveled by the neutron until its energy is decreased below the ${}^{39}\text{K}(n, p)$ threshold, $\bar{\sigma}_{39}$ is the average value of ${}^{39}\text{K}(n, p)$ cross section, and N_{39} is the atom density of ${}^{39}\text{K}$ nuclide. Using approximate values,

$$P_{39}(5 \text{ MeV}) \approx 40 \text{ cm} \times 0.3 \text{ barn} \times 9 \times 10^{20} \text{ atoms cm}^{-3} = 0.01 \quad (23)$$

- We use MCNP6 code for a more educated calculation
(Monte Carlo N-Particle, <http://mcnp.lanl.gov>)



Results

Table 5: Production rates of ${}^4\text{He}$, neutrons, ${}^{21}\text{Ne}$, ${}^{39}\text{Ar}$.

| Composition | Production rates in atoms/kg-yr | | | |
|-------------------------------|---------------------------------|----------|--------------------|--------------------|
| | ${}^4\text{He}$ | neutrons | ${}^{21}\text{Ne}$ | ${}^{39}\text{Ar}$ |
| <i>RG03</i> , Upper Crust | 1.64×10^{10} | 10350 | 0.151 | 29.9 |
| <i>RG03</i> , Middle Crust | 8.98×10^9 | 6232 | 0.159 | 14.7 |
| <i>RG03</i> , Lower Crust | 1.53×10^9 | 1156 | 0.103 | 0.792 |
| <i>SS04</i> , Depleted Mantle | 2.51×10^7 | 22.2 | 0.0380 | 2.58E-04 |

Table 6: Production rates of ${}^{21}\text{Ne}$ by (α, n) and (n, α) and ${}^{21}\text{Ne}/{}^4\text{He}$ ratio.

| Composition | ${}^{21}\text{Ne}$ prod. rate in atoms/kg-yr | | | % contrib. | | ${}^{21}\text{Ne}/{}^4\text{He}$ |
|-------------------------------|--|---------------|-------|---------------|---------------|----------------------------------|
| | (α, n) | (n, α) | Total | (α, n) | (n, α) | |
| <i>RG03</i> , Upper Crust | 863.8 | 0.15 | 864 | 99.98 | 0.02 | 5.26×10^{-8} |
| <i>RG03</i> , Middle Crust | 474.7 | 0.16 | 475 | 99.97 | 0.03 | 5.29×10^{-8} |
| <i>RG03</i> , Lower Crust | 79.3 | 0.10 | 79.4 | 99.9 | 0.1 | 5.18×10^{-8} |
| <i>SS04</i> , Depleted Mantle | 1.18 | 0.038 | 1.22 | 96.9 | 3.1 | 4.85×10^{-8} |

Yokochi et al. (1997): ${}^{21}\text{Ne}/{}^4\text{He} = 4.5 \times 10^{-8}$

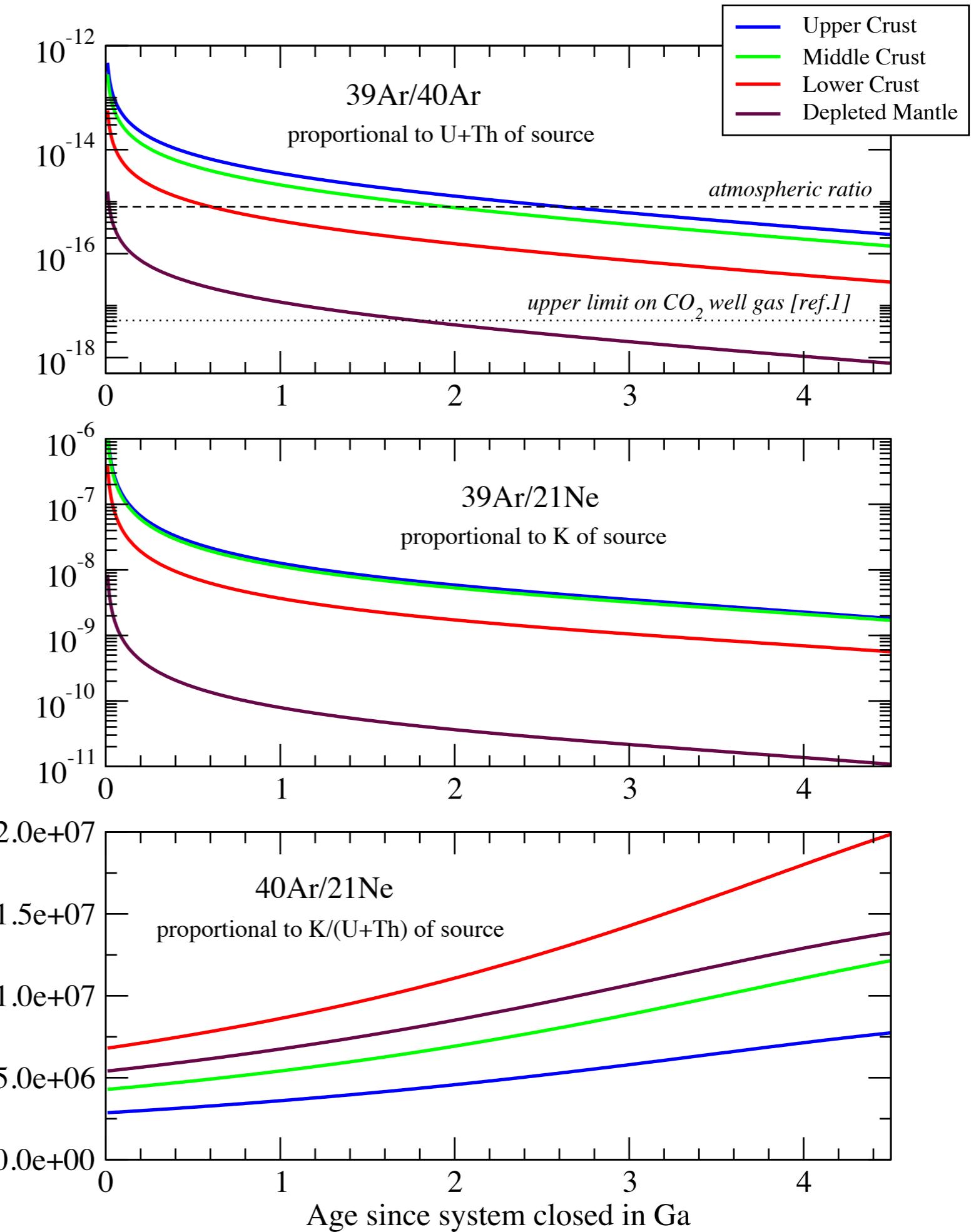
Isotopic ratios

Trade-off between composition and “age” of system

Closed system:

^{21}Ne , ^{40}Ar accumulate

^{39}Ar steady-state level



Summary

- New calculations of nucleonic production rates
- Combination of state-of-the-art nuclear physics tools (TALYS, MCNP6) and my own code to put everything together
- (Re-)Learned aspects of nuclear physics
- Results of interest to hydrology, dark matter experiments, crustal & mantle geochemistry
- We expect the ^{39}Ar detection methods to improve